

The North Pyrenean Frontal Thrust: structure, timing and late fluid circulation inferred from seismic and thermal-geochemical analyses of well data

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Abstract – During orogenesis, large-scale thrusts as orogenic fronts can act as conduits and/or barriers for fluid flow. Unravelling the timing and modes of tectonic activation of large-scale faults is crucial to understanding the relationship between fluid flow and deformation. The North Pyrenean Frontal Thrust (NPFT) corresponds to a major basement-involved thrust responsible for the northward overthrust of the pre-orogenic sediments on top of the Aquitaine Foreland Basin. This study questions the timing of activation of this thrust, its geometry, the nature of the last fluids, which circulated there, and its role on the circulation of fluids. The structural study confronted to published thermochronology data led to determine the timing of the two tectonic activations during the NPFT compression phase and to relate them to the fluid circulations. We constrain the first activation at Campanian times and link it to the leak of the deep gas reservoir present in depth, as the NPFT acted as a conduit. Then the NPFT acted as a barrier, probably due to the breccia consolidation during the Paleocene quiescence period. Finally, the Eocene-Oligocene reactivation led to fluid circulation of high salinity fluids from the Triassic evaporites leaching. This latter event is associated with a fracturing event and the late generation of calcite veins studied here. This is the first study in the Pyrenees directly applied to the NPFT which uses the association between fluid inclusions study, seismic and thermochronological data. It highlights that the NPFT may be an important structure responsible of the leakage of deep hydrocarbons reservoirs. It also shows the importance of the determination of the activation steps of large-scale faults to decipher the origin of fluid circulations in space and time.

Keywords: Pyrenees / North Pyrenean Frontal Thrust / fluid inclusions / thermochronology / stable isotopes / evaporites

Résumé – Le chevauchement frontal Nord Pyrénéen : structure, timing et circulation de fluide tardif déduits d'analyses sismiques et thermo-géochimiques de données de puits. Au cours d'une orogénèse, les chevauchements majeurs en tant que fronts orogéniques peuvent agir comme conduits et/ou barrières pour la circulation de fluides. Il est essentiel de comprendre la chronologie et le mode d'activation de telles failles à grande échelle, pour comprendre la relation entre les circulations fluides et la déformation. Le Chevauchement Frontal Nord Pyrénéen (CFNP) correspond à un chevauchement majeur ancré dans le socle et responsable du chevauchement vers le nord des sédiments pré-orogéniques au sommet du Bassin d'avant-pays Aquitaine. Cette étude interroge le timing d'activation de ce chevauchement, sa géométrie, la nature des derniers fluides qui y ont circulé et son rôle dans la mise en circulation de ces fluides. L'étude structurale confrontée à des données de thermochronologie déjà publiées, a conduit à déterminer la chronologie des

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deux activations tectoniques de la phase de compression du CFNP et à les relier aux circulations fluides. Nous contrainsons la première activation au Campanien et a probablement conduit à la fuite du réservoir de gaz présent en profondeur, le CFNP jouant le rôle de conduit. Ensuite, il a agi comme une barrière, probablement en raison de la consolidation de la brèche pendant la période de quiescence du Paléocène. Enfin, la réactivation Eocène-Oligocène a conduit à une circulation de fluides très salés issus du lessivage des évaporites du Trias. Ce dernier événement est associé à un événement de fracturation et à la génération tardive des veines de calcite étudiées ici. Il s'agit de la première étude dans les Pyrénées directement appliquée au NPFT qui utilise une association entre étude des inclusions fluides et données sismiques et thermochronologiques. Elle met en évidence que le NPFT peut être une structure importante responsable de la fuite des réservoirs d'hydrocarbures profonds. Elle montre également l'importance de la détermination des étapes d'activation des chevauchements de grande échelle pour déterminer l'origine des circulations fluides dans l'espace et le temps.

Mots clés : Pyrénées / chevauchement frontal Nord Pyrénéen / inclusions fluides / thermochronologie / isotopes stables / évaporites

1 Introduction

Large-scale faults at orogenic fronts are privileged places to study the relationships between long-term deformation processes with short-term episodic fluid circulations. Fluid/rock interactions studies in faults zones were originally driven by industrial needs to understand the role and timing of the fluid circulations in these faults at the origin of oil and gas and ore deposits (*e.g.*, Travé *et al.*, 1997, 1998; Boles *et al.*, 2004; Eichhubl *et al.*, 2009; Weis *et al.*, 2012; Grare *et al.*, 2018). Fluid flow associated to active faults is governed by different mechanisms such as fault-valve action, seismic pumping, anisotropic permeability, or hydraulic fracturing by increasing pore pressure (*e.g.*, Sibson, 1981, 2000; McCaig, 1988; Magee and Zoback, 1993; Bense and Person, 2006). Foreland thrusts are known to usually act as conduits for fluids (Travé *et al.*, 2000; Cruset *et al.*, 2018; Smeraglia *et al.*, 2020) driving circulation of different types of fluids dependent on their tectonic history and sources (*e.g.*, Travé *et al.*, 2007; Lacroix *et al.*, 2014). Only recently, the increasing ability of geothermochronometric techniques to resolve the timing of complex fluid flow events open the way to significantly improve our knowledge of rock-fluids interactions (Louis *et al.*, 2019; Milesi *et al.*, 2019, 2020). Source, nature and timing of the fluid circulations in the central part of the South Pyrenean Zone are well documented for Cenozoic syn-orogenic deformation phases (*e.g.*, Rye and Bradbury, 1988; Travé *et al.*, 2000; Lacroix *et al.*, 2014; Beaudoin *et al.*, 2015; Crognier *et al.*, 2018; Cruset *et al.*, 2018). In the North Pyrenean Zone and Aquitaine Basin most studies dealing with fluid circulations have focused on the Mesozoic syn-rift history (Salardon *et al.*, 2017; Corre *et al.*, 2018; Quesnel *et al.*, 2019; Renard *et al.*, 2019; Elias-Bahnan *et al.*, 2020; Motte *et al.*, 2021; Cathelineau *et al.*, 2021), the syn-orogenic imprint remaining poorly constrained. One of the main syn-orogenic faults carrying fluids in the north of Pyrenees is the North Pyrenean Frontal Thrust (NPFT; Connan and Lacrampe-Couloume, 1993) which separate the Aquitaine Foreland Basin from the North Pyrenean Zone (NPZ; Fig. 1).

The NPFT is a large basement-involved structure responsible for the northward overthrust bringing the Early Cretaceous syn-rift successions (Barremian-Albian) on top of the late Cretaceous to Paleogene deposits of the Aquitaine

Basin. While studies on the tectonic evolution of the NPFT have been carried out in the Central and Eastern Pyrenees (Mattauer and Proust, 1962; Charrière and Durand-Delga, 2004; Ford *et al.*, 2016; Labaume and Teixell, 2020), the arguments used to constrain its activity to the west remain very scarce as the western fault trace is buried below a thin layer of Miocene post-orogenic sediments. The role and timing of the NPFT in controlling fluid circulation has not yet been demonstrated.

In this study, we couple new geochemical and interpretation of seismic reflection lines with recently published thermochronological analyses to determine the timing of thrust activity and the nature and origin of the late fluid circulations in the NPFT. The fluid circulation model is built in accordance with previous more detailed geochemical studies. This contribution integrates results from both Fluids and Orogen R&D Total projects (please refer to the Website for more information on these projects: <https://convergent-margins.com/>) to document fluid-rock interactions associated with large-scale convergent complex structures.

2 Main deformation phases in the Pyrenean orogen

The Pyrenees are an alpine-type mountain belt formed between the Santonian and the Miocene as a result of the convergence between Iberia and Europe. Collision involves a Variscan basement and a Mesozoic sedimentary cover that underwent Early Cretaceous rifting. Evidence for the Early Cretaceous rifting are best preserved in the North Pyrenean Zone (NPZ; Fig. 1), where the crustal thinning led to hyper-extension and mantle exhumation (*e.g.*, Fabriès *et al.*, 1991, 1998; Lagabrielle and Bodinier, 2008; Jammes *et al.*, 2009; Lagabrielle *et al.*, 2010; Clerc *et al.*, 2012, 2013; Masini *et al.*, 2014; Tugend *et al.*, 2014). The peculiar thermal evolution of the Pyrenean rifting is well documented and has revealed widespread HT-LP metamorphism (Ravier, 1959; Azambre and Rossy, 1976; Bernus-Maury, 1984; Dauteuil and Ricou, 1989; Golberg and Leyreloup, 1990; Clerc and Lagabrielle, 2014; Clerc *et al.*, 2015; Lescoutre *et al.*, 2019; Ducoux *et al.*, 2019, 2021b), and the establishment of high geothermal gradient (up to 80 °C/km; Vacherat *et al.*, 2014;

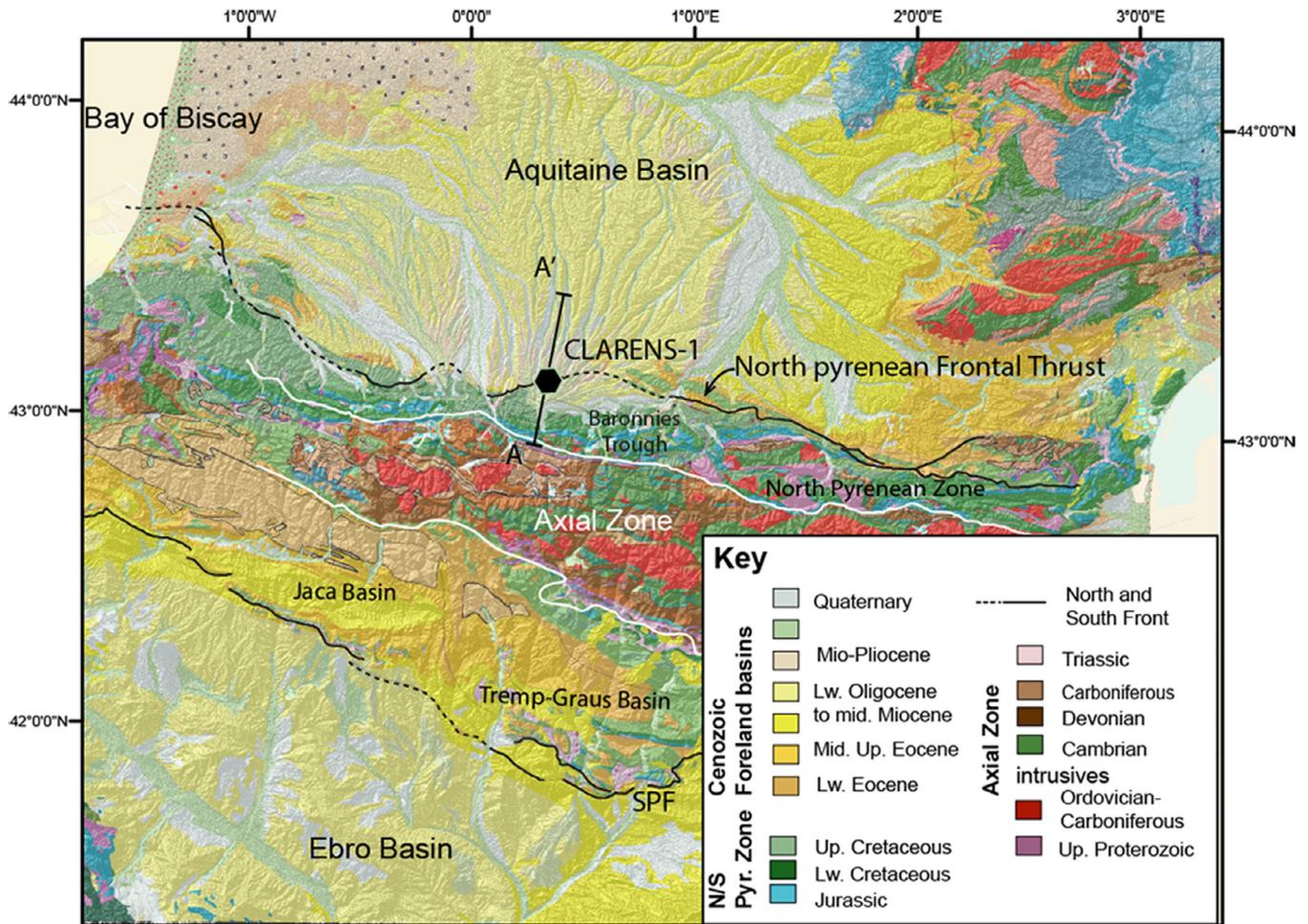


Fig. 1. Geological map of the Pyrenean Range with location of the main features discussed in this study. AA' is the section detailed in [Figure 3](#).

[Renard *et al.*, 2019](#)) in Early Cretaceous sediments of the Mauléon Basin, in the Western Pyrenees.

The onset of Iberia/Europe convergence is documented by the oldest sea-floor spreading magnetic anomalies in the Atlantic Ocean dated to the Santonian (*e.g.*, [Roest and Srivastava, 1991](#); [Olivet, 1996](#); [Rosenbaum *et al.*, 2002](#); [Sibuet *et al.*, 2004](#); [Macchiavelli *et al.*, 2017](#)) and geological evidence in the Pyrenees of a late Santonian unconformity and deposition of clastic sediments ([Garrido *et al.*, 1972](#); [Muñoz, 1992](#); [Vergés *et al.*, 1995](#); [Teixell, 1998](#); [Vergés and García Senz, 2001](#); [García Senz, 2002](#); [McClay *et al.*, 2004](#); [Mouthereau *et al.*, 2014](#); [Labaume and Teixell, 2020](#)). This early convergence stage is marked by the underthrusting of the exhumed mantle and hyperextended rift domain ([Mouthereau *et al.*, 2014](#); [Ford *et al.*, 2016](#); [Teixell *et al.*, 2016](#); [Gómez-Romeu *et al.*, 2019](#)). As deformation propagated southward by the accretion of more proximal domains of the Iberian margin, the inversion of former rift structures such as the Boixols fault occurred, driving deposition of Campanian-Maastrichtian flysch (see synthesis in [Martin-Chivelet *et al.*, 2019](#)). This early orogenic phase is supported by substantial evidence arguing for an early phase of cooling at about 70 Ma based on low-temperature thermochronology (*e.g.*, [Whitchurch *et al.*, 2011](#); [Beamud *et al.*, 2011](#); [Mouthereau *et al.*, 2014](#); [Ternois *et al.*, 2019](#); [Waldner *et al.*, 2021](#)).

Farther west, marine geophysical constraints further document Santonian-Campanian heterogeneous reactivation of the northern Bay of Biscay ([Thinon *et al.*, 2001](#)). This is also suggested in the Basque-Cantabrian Basin by the transition from calcareous to siliciclastic flyschs in the early Campanian times. Farther north, subsidence analyses and flexure modelling reveal that the onset of foreland deflection of the European plate occurred in the late Cretaceous in response to incipient orogenic loading ([Vergés *et al.*, 2002](#); [Ford *et al.*, 2016](#); [Angrand *et al.*, 2018](#)).

The mature collisional stage marked by widespread orogenic exhumation in the Pyrenees occurred from ~50 Ma to ~20 Ma ([Muñoz, 1992](#); [Mouthereau *et al.*, 2014](#) and references therein) suggests an onset of accretion of necking domains as argued by modelling ([Jourdon *et al.*, 2019](#)). This interval defines a period of wedge widening that is reflected in the southern fold-and-thrust belt by thin-skinned deformation accommodated in the Triassic evaporitic level (*e.g.* Tremp-Graus Basin, see [Fig. 1](#)). In the North Pyrenean Zone, however, thick-skinned or basement-involved inversion is the dominant mode of deformation. This is associated with less mechanical decoupling between the basement and supra-salt sedimentary units that indicate comparatively less shortening on the retro-wedge side of the orogen ([Jourdon *et al.*, 2019](#)). In the Axial Zone, deformation is characterized

by crustal nappe stacking. The positive feedback between tectonics and climate resulted in intense exhumation of the Axial Zone between ~ 40 Ma and ~ 30 Ma at rates up to 2 km/Myr (Fitzgerald *et al.*, 1999; Fillon and van der Beek, 2012). Contraction and flexural subsidence associated with topographic loading appear to have ceased slightly earlier (~ 23 Ma; Rocher *et al.*, 2000; Ortiz *et al.*, 2020) in the northern retro-foreland than in the southern pro-foreland (~ 18 Ma, Oliva-Urcia *et al.*, 2019).

3 Borehole description

Throughout this study, we use the Clarens-1 borehole (CLS-1) which was drilled in the 60's at the early stages of hydrocarbon exploration in the Aquitaine Basin by Elf and is now part of the patrimonial TOTAL dataset. This borehole is located in the retro-foreland of the Pyrenean belt, at the apex of the Lannemezan plateau. The borehole penetrates the Cenomanian series made of alternation of marls with a few sandstone beds from 200 m to ~ 1000 m depth, then a disturbed ~ 500 m thick zone consisting of polygenic breccias into the Cenomanian marls, interpreted as the North Pyrenean Frontal Thrust (NPFT) damage zone. Finally, Maastrichtian and Campanian rocks have been drilled between ~ 1500 m to 3500 m (Fig. 2). Present-day temperature measured in the Clarens-1 borehole indicates a value of 65 °C at 2698 m (geothermal gradient of 24 °C/km), consistent with the temperatures measured in other boreholes nearby (TOTAL internal data).

4 Methodology

4.1 Seismic interpretation

The structure of the NPFT is inferred from seismic reflection data and boreholes information available from the TOTAL patrimonial dataset. In order to detail the architecture of the NPFT and the adjacent Aquitaine Basin, a new geological cross-section is proposed based on the interpretation of a N-S seismic profile (85SAV2, courtesy of Total SE) and 7 wells (Lannemezan-1 & 2, Clarens-1, 2 & 102, Monlong-1 and Recurt-1) (Fig. 3).

4.2 Petrography

Rock samples from the breccia zones at 1000, 1487 and 1489 m in the Cenomanian formation, were collected to determine the different generations of cements and veins. Observations were made under an optical microscope and systematically supplemented by cathodoluminescence petrography performed with a NewTec Scientific cathodyne (10–12 kV and 100 mA) coupled to a Nikon eclipse 80i microscope and a digital iDS camera.

4.3 Microthermometric analysis

Samples for microthermometry analysis were taken from the Clarens-1 borehole. First, we conducted macroscopic observations over the entire available core of the Clarens-1 borehole to characterize the structure and composition of late calcite veins (Cal 3) used for fluid inclusions study (see Sect. 5.1 for the

relative timing of the different generations). Due to limited sampling possibilities, only veins in the fault zone at 1000 m and 1487 m depth were used for fluid inclusions analysis. The host-rock is a polygenic breccia within the Cenomanian marls crosscut by calcite veins where fluid inclusions were analyzed (Figs. 4A and 4B). A total of 42 fluid inclusions were studied on double polished 150 µm-thick sections from the calcite veins. Microthermometric measurements were performed by [®]HES Expertise Services, on a Linkam MDS 600 heating-cooling stage connected to a Sony Exwave HAD3 color video camera mounted on an Olympus BX 51 microscope at GeoRessources, Vandoeuvre-lès-Nancy. The microthermometric stage was thermally calibrated using CO₂ synthetic fluid inclusions standards for $T_{m(\text{CO}_2)}$ (melting temperature of CO₂), H₂O synthetic fluid inclusions standards for $T_{m(\text{ice})}$ (melting temperature of ice), and cross-calibrated natural alpine fluid inclusions (“home” standard) for T_h (homogenization temperature). The accuracy of measurements is estimated at ± 0.2 °C for $T_{m(\text{ice})}$ and ± 0.5 °C for T_h . In order to detect the presence of oil-rich fluid inclusions, epi-UV fluorescence observations were conducted using a Zeiss AxioImager.A1m microscope with an excitation filter centered at 365 ± 5 nm and an emission filter long pass around 400 nm.

4.4 Vitrinite reflectance analysis

Only 4 organic matter-rich samples suitable for reflected light microscopy were found on the Clarens-1 core. They were crushed between 1 to 2 mm for observation of the organic facies and for estimation of the fluorescence intensity of the groundmass. A concentrate of organic matter using heavy liquid was used for random measurements. Random reflectance measurements were conducted using a LEICA DM4P or DM4500P microscopes, with objectives of x20 and x50. Calibration was made with three standards method: Saphir (0.586%), Glass (1.25%) for low maturity, Glass (1.25%) and Zirkonia (3.115%) for medium maturity, and Zirkonia (3.115% to 7.32%) for high maturity.

5 Results

5.1 Structure of the North Pyrenean Frontal Thrust

The cross-section we investigate stretches between the Baronnies Trough to the south and the Aquitaine Basin to the north (Fig. 3). The Baronnies Trough corresponds to a syn-rift basin that developed over a hyperthinned crust and an exhumed mantle (*e.g.*, Espurt *et al.*, 2019). As determined in the Lannemezan-1 well, the northern part of Baronnies Trough located on the hanging wall of the NPFT is filled by Albo-Cenomanian syn-rift successions which sit on the Upper Triassic evaporites (Fig. 3). It is noteworthy the lack of the pre-rift sequence in the northern part of the Baronnies Trough. Indeed, this was not recognized in the Lannemezan-1 & 2 wells where the syn-rift deposits overlay the Upper Triassic evaporites. Below this non-isopachous evaporitic sequence, the top of the basement can be highlighted by high-amplitude reflectors (Fig. 3). Large accumulations of salt-bearing rocks (> 1.5 km) are documented by Lannemezan-1 & 2 wells along the NPFT. The polygenic breccia, which are

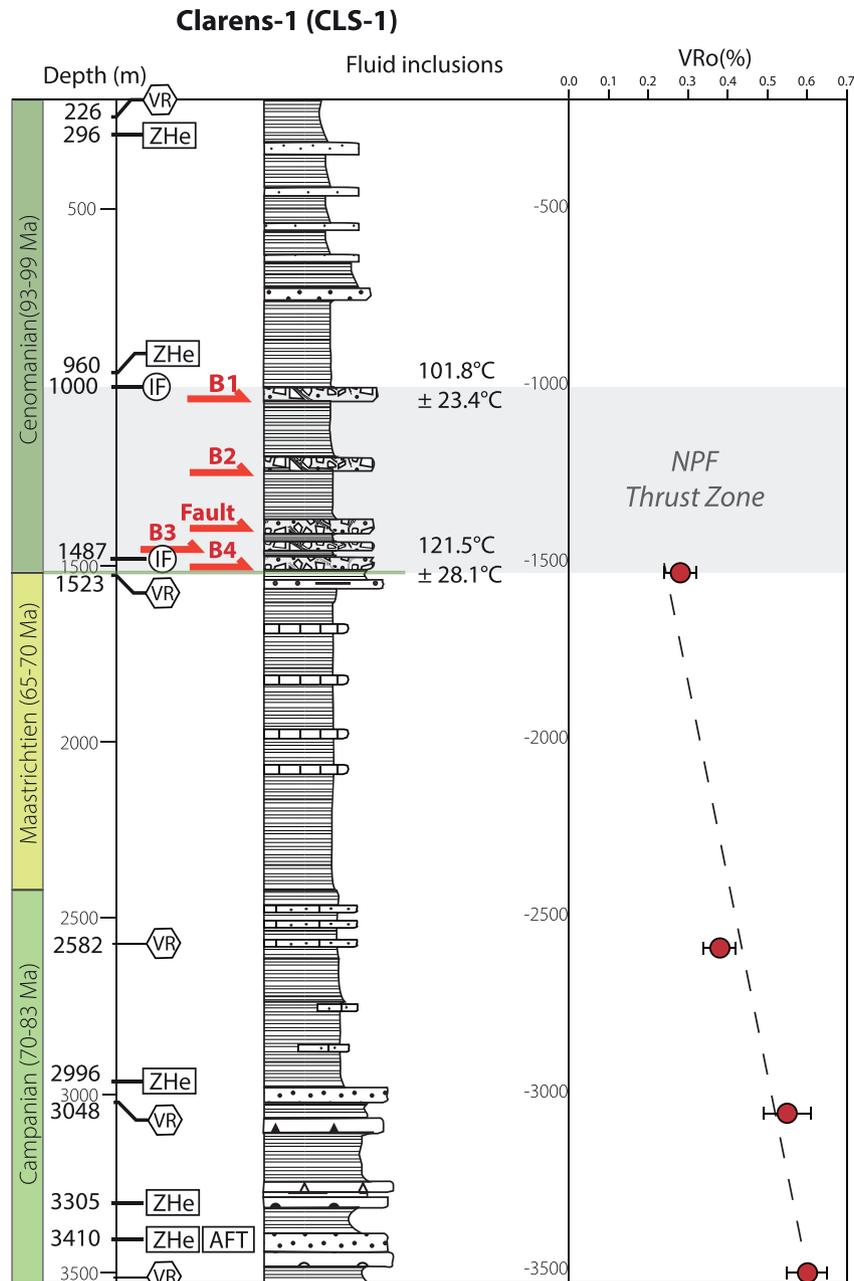


Fig. 2. Schematic log of the Clarens-1 (CLS-1) borehole with the vitrinite reflectance data used in this study. Note the 500 m large zone which presents the fault zone and the different levels of sedimentary breccia (B1 to B4) observed on the drill core. VR, AFT, ZHe and IF refer to the position of the Vitrinite Reflectance data, Apatite Fission Track sample, (U-Th)/He on zircons samples and fluid inclusions samples, respectively.

described below, were drilled in Lannemezan-2, Clarens-1, 2 & 102 wells and are located along the south-dipping NPFT.

North of the NPFT, in the Aquitaine Basin, the top-basement reflector is observed at 3 s TWT, indicating a thick (~7.5 km) accumulation of sediments in the footwall. The NPFT is covered by Miocene and younger sediments. Below, the Eocene deposits show growth-stratal patterns defined by beds thickening towards the NPFT and wedging out to the frontal fold drilled at Recurt-1 (Fig. 3). By contrast, the Paleocene beds are isopachous. The thick Late Cretaceous (Santonian to Maastrichtian) accumulations that are inferred to

mark the early stage of convergence show growth strata wedging toward the NPFT. These syn-convergence series further depict a large tight recumbent fold along the NPFT as documented by reverse dip measured in Lannemezan-1 & 2 wells (Fig. 3). Although poorly dated in detail, the growth strata geometries suggest two phases of deformation associated with motion along the NPFT during the Santonian-Maastrichtian and in the Eocene times.

The bottom of the early syn-convergence sequences displays several apparent downlaps and show they were deposited unconformably above the post-rift sequence (Fig. 3).

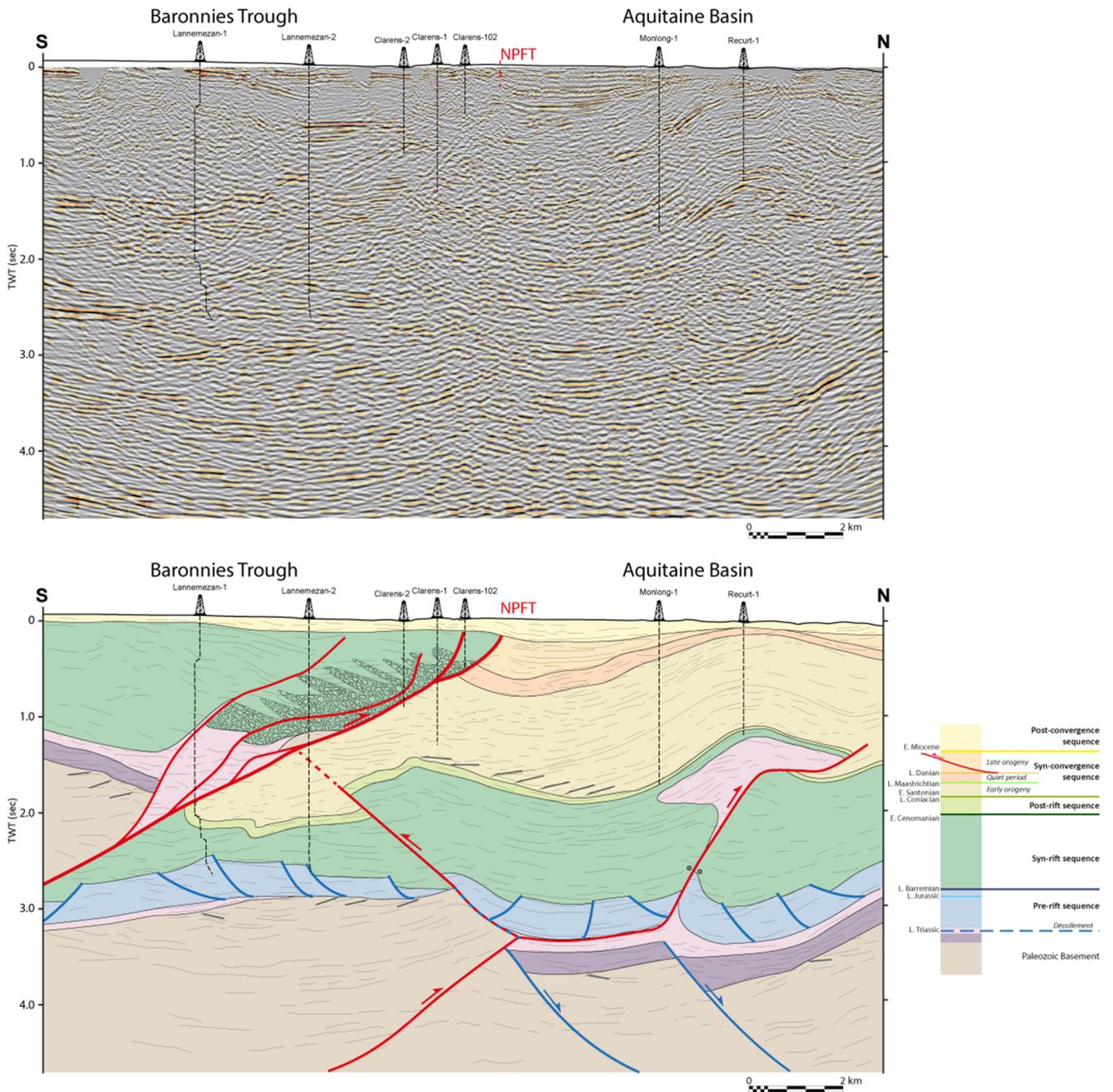


Fig. 3. Structure of the North Pyrenean Frontal Thrust along N-S profile located in the central Pyrenees (see Fig. 1 for location). (a) N-S Seismic reflection profile (85SAV2, courtesy of Total SE) with the position of Lannemezan-1 & 2, Clarens-1, 2 & 102, Monlong-1 and Recurt-1. (b) Section of the interpreted seismic profile and calibrated with wells. This section shows the geometry of the North Pyrenean Frontal Thrust, and the structure of the relatively preserved Aquitaine Basin.

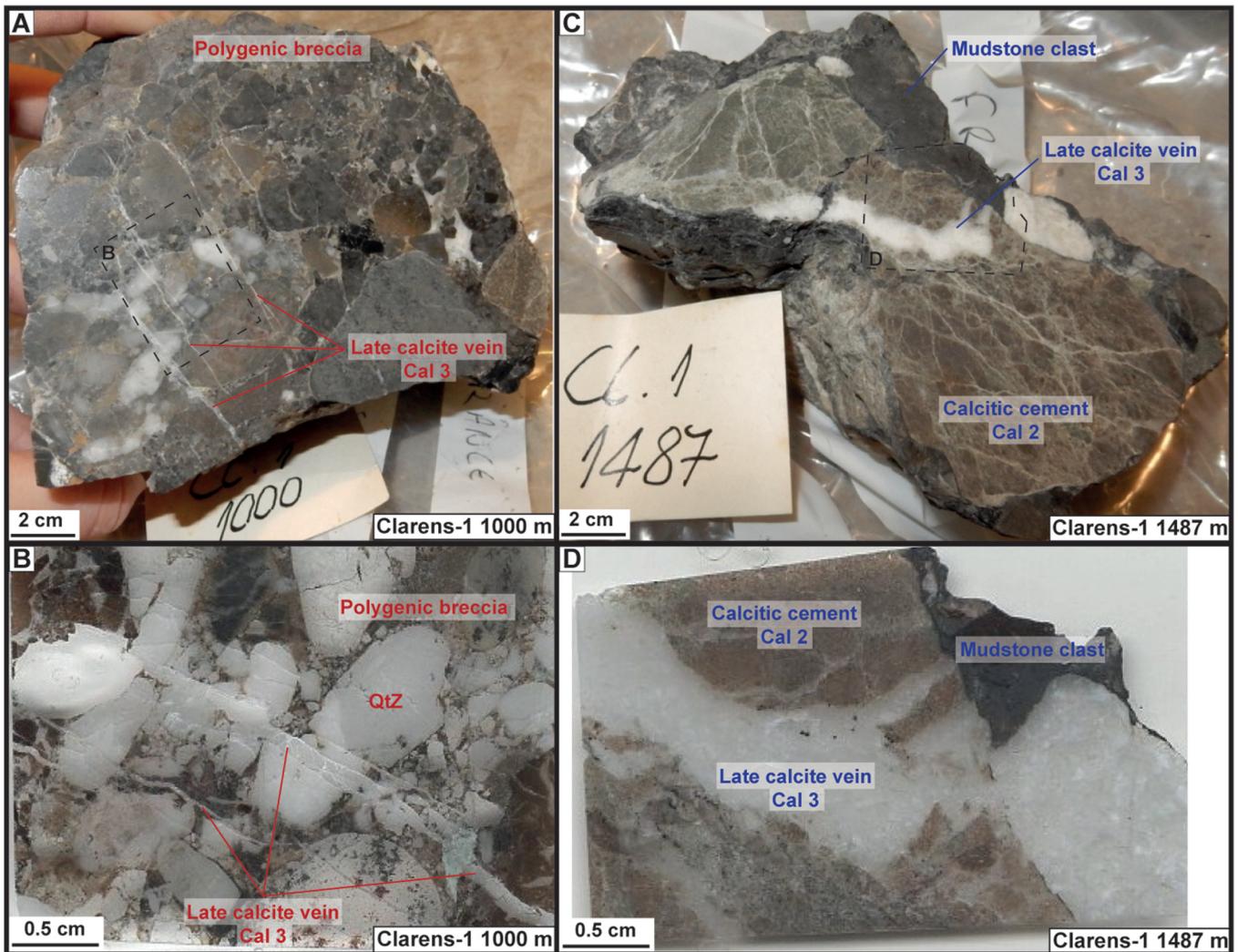


Fig. 4. Photographs of macroscopic samples and thin sections from Clarens-1 at 1000 m (A and B) and 1487 m (C and D) depth, showing the polygenic breccia of the NPF Thrust Zone crosscut by late calcite veins (Cal 3) used for fluid inclusions study.

They indicate that the underlying former rift template is sealed and relatively preserved from the collisional imprint.

The thickness of the post-rift sequence although calibrated in Monlong-1 et Rencurt-1 wells only appears to be isopachous in the footwall of the NPFT. The thickness of the underlying syn-rift sequence is determined in contrast by Lannemezan-1 & 2 in the south, as well as by Monlong-1 and Recurt-1 in the north. These deposits show a constant thickness within the Aquitaine Basin. However, their thickness varies to the south as revealed by thinning toward the NPFT and where they are cross-cut by a welded salt diapir composed of the Upper Triassic evaporites (Fig. 3). The pre-rift sequence is discontinuous, anisopachous and is affected by normal faulting. In the northern part of the section, the pre-rift series overlying the Upper Triassic evaporites shows two synclines on both sides of a salt diapir. Underneath the Baronnie Trough and the NPFT, the pre-rift sequence becomes shallower and sits on the basement.

The top-basement geometry can be identified because of the high-amplitude reflectors and the top-lap, corresponding to the angular unconformity between the Lower Triassic deposits

and the Variscan basement. Furthermore, As the pre-rift sediments (including Upper Triassic evaporites and Jurassic carbonates) accommodate the geometry of the basement (*e.g.* shift) their shape highlights the top-basement geometry. Basement-involved tilted blocks bounded by north-dipping normal faults are observed in the northern part of the section. Below the Clarens-1, the top-basement geometry shows a shift of ~ 1 s TWT (also outlined by Permian-Triassic deposits), which correlates with the stretching recorded in the pre-rift cover and thick accumulations of syn-rift sediment.

This shift of the top-basement indicates the north-dipping normal fault was affected by subsequent deformation related to the collision event. This type of architecture can be observed further west in the neighboring Arzacq Basin (Ducoux *et al.*, 2021a). The upper part of the tilted block is short-cut by a reverse thick-skinned fault forming a fish-tail structure in the Aquitaine Basin (Fig. 3). Part of the thick-skinned deformation was accommodated in the sedimentary cover by a back-thrust responsible for the shift of pre-, syn-, post-rift and syn-orogenic sediments toward the NPFT. The downlaps of the syn-orogenic sediments encapsulated the open fold related to

this back-thrust. Below Recurt-1 and Monlong-1 wells, the displacement along the thick-skinned basement thrust ramp is seen to be accommodated in the weak Upper Triassic salt layer. The fault ramp then connects to the welded diapiric structure that affects the syn-rift sediments. This deformation is eventually accommodated in gentle folding of the syn-orogenic sediments (Fig. 3).

5.2 Breccia and calcite vein petrography

The core samples collected in Clarens-1 well across the NPFT fault zone are made of an alternation of breccias and marls. Breccias are encountered in the Cenomanian formation at 1000 m, 1241 m, 1388 m, 1442 m and 1487 m (Fig. 2). They are polygenic, containing clasts of sandstones, sub-angular quartz, dolomites, and fine-grained black siliciclastic argillites, cemented by calcite (Fig. 4). The different clast lithologies and the absence of cataclastic deformation argue for the sedimentary origin of the breccias. Such clasts probably come from the various sedimentary formations present at depth or have been emplaced during episodes of salt tectonics. Cataclastic deformation is encountered in marls and breccias between 1388 m and 1399 m. They are made of sheared and striated marls surrounding clasts corresponding to the sedimentary breccias described above.

Four generations of calcite (as cements and veins) are observed at 1000 m (Fig. 5). The first one (Cal 0; Figs. 5B–5D) corresponds to a fracture formed in the sedimentary rock before its brecciation. The breccia is cemented by two non-luminescent calcites (Cal 1 and Cal 2, the first being the least luminescent; Figs. 5B–5D) which are anterior to a luminescent calcite (Cal 3; Figs. 5B–5D). The breccias inside the NPFT fault zone is crosscut by calcite veins. It is noteworthy that the calcite veins are only observed in the breccias and not in the host Cenomanian marls. At 1489 m, two generations of fractures are filled by similar luminescent calcite than Cal 3 observed at 1000 m (Fig. 5F).

5.3 Fluid inclusions analysis

Fluid inclusions (FIs) found in our two samples of Cal 3 are biphasic with the aqueous phase volumetrically dominant over the vapor phase (Fig. 6). They all homogenized into liquid phase. No oil-rich fluid inclusions were observed. In the two veins, FIs display similar salinity around 20.9 ± 3.4 wt% eq. NaCl. At 1000 m, T_h values ranged between 63.1°C and 145.1°C with a mean at $101.8 \pm 23.4^\circ\text{C}$; at 1487 m they show slightly higher T_h with a range between 72.2°C and 174.7°C and a mean at $121.5 \pm 28.1^\circ\text{C}$ (Fig. 7; Tab. 1). Note that the observed minimum T_h (63.1°C) is higher than the present-day temperature in the borehole (around 36°C at 1500 m; see Sect. 3), implying no T_h reset here. All the fluid inclusions analyzed are observed in intragranular planes in the calcite (Fig. 6), which indicates primary fluid inclusions (Van den Kerkhof and Hein, 2001). A large range of T_h values are measured for the same fluid inclusions plan, indicating that no petrographic feature can explain such a large range. For instance, inclusions 11 and 12 at 1000 m depth are located in the same plan (Fig. 6C) but show T_h at 109.0°C and 66.9°C

respectively. Another example corresponds to inclusions 5 and 6 obtained at 1487 m depth (Fig. 6D) where T_h at 174.7°C and 72.2°C are measured. Eutectic temperatures (T_e) depend of the nature of the salt dissolved in the fluid inclusion. T_e between -37.9°C and -45.8°C with a mean at $-41.7 \pm 2.9^\circ\text{C}$ were determined at the 1000 m level (Tab. 1). Such T_e corresponds to a mixing between NaCl–CaCl₂–H₂O and NaCl–MgCl₂–H₂O systems which have T_e at -37°C and -52°C , respectively (Davis *et al.*, 1990). A single T_e at -64.5°C could be determined for the 1487 m level (Tab. 1) indicating a possible addition of LiCl or exhibit metastability (Cathelineau *et al.*, 1994; Dubois *et al.*, 2010).

5.4 Vitrinite reflectance

Vitrinite reflectance analysis reveals % R_o values increasing with depth from $0.28 \pm 0.04\%$ in the upper part of the core samples (1523 m) to $0.60 \pm 0.05\%$ in the lower part (3501 m; Tab. 2). These two estimates indicate that these samples are made of sediments with a low thermal maturity. Such low reflectance values in samples located in the footwall of the NPFT suggest a normal burial-dependant trend and equilibration under a gradient of $30^\circ\text{C}/\text{km}$.

6 Discussion

6.1 Tectonic evolution

6.1.1 From syn-rift to post-rift evolution

The increase of syn-rift sediment thickness in the hanging wall of the NPFT (Baronnies Trough) indicates the NPFT formed by the reactivation of a former south-dipping normal fault, bounding a hyperextended domain with mantle exhumation (Fig. 8A). Raman Spectroscopy on Carbonaceous Materials performed on sedimentary rocks of the Baronnies Trough further reveals that this basin experienced temperatures up to 480°C (Espurt *et al.*, 2019; Ducoux *et al.*, 2021b) indicating high thermal gradient achieved during hyperextension. The polygenic breccias observed in the Clarens-1 boreholes (Figs. 4A and 4B) are inferred to be proximal clastic sediments deposited in the hanging wall of a south-dipping normal fault (Fig. 8A). The thick accumulations of Upper Triassic evaporites documented by the Lannemezan-1 well, suggest that a large diapiric structure may have formed associated with this large-scale normal fault (Fig. 8A). The paleo-high located between the Baronnies Trough and the Aquitaine Basin is highlighted by the preservation of the pre-rift sequence and thinner syn-rift deposit sealed by a thin post-rift deposit.

To the North, the vitrinite reflectance values shows normal depth-dependant burial in the footwall of the NPFT suggesting the rift-related thermal imprint was relatively modest (Fig. 2; Tab. 2). The northern margin of the basement high is bounded by a succession of high-angle normal faults shaping a succession of tilted blocks. The pre-rift sequence was detached along the Upper Triassic evaporites and folded during the ascent of a huge diapiric structure recognized in the Recurt-1 well. The spatial distribution of salt diapirs appears related to the location of the main basement-involved faults. During rifting, the extension and thinning was more important to the

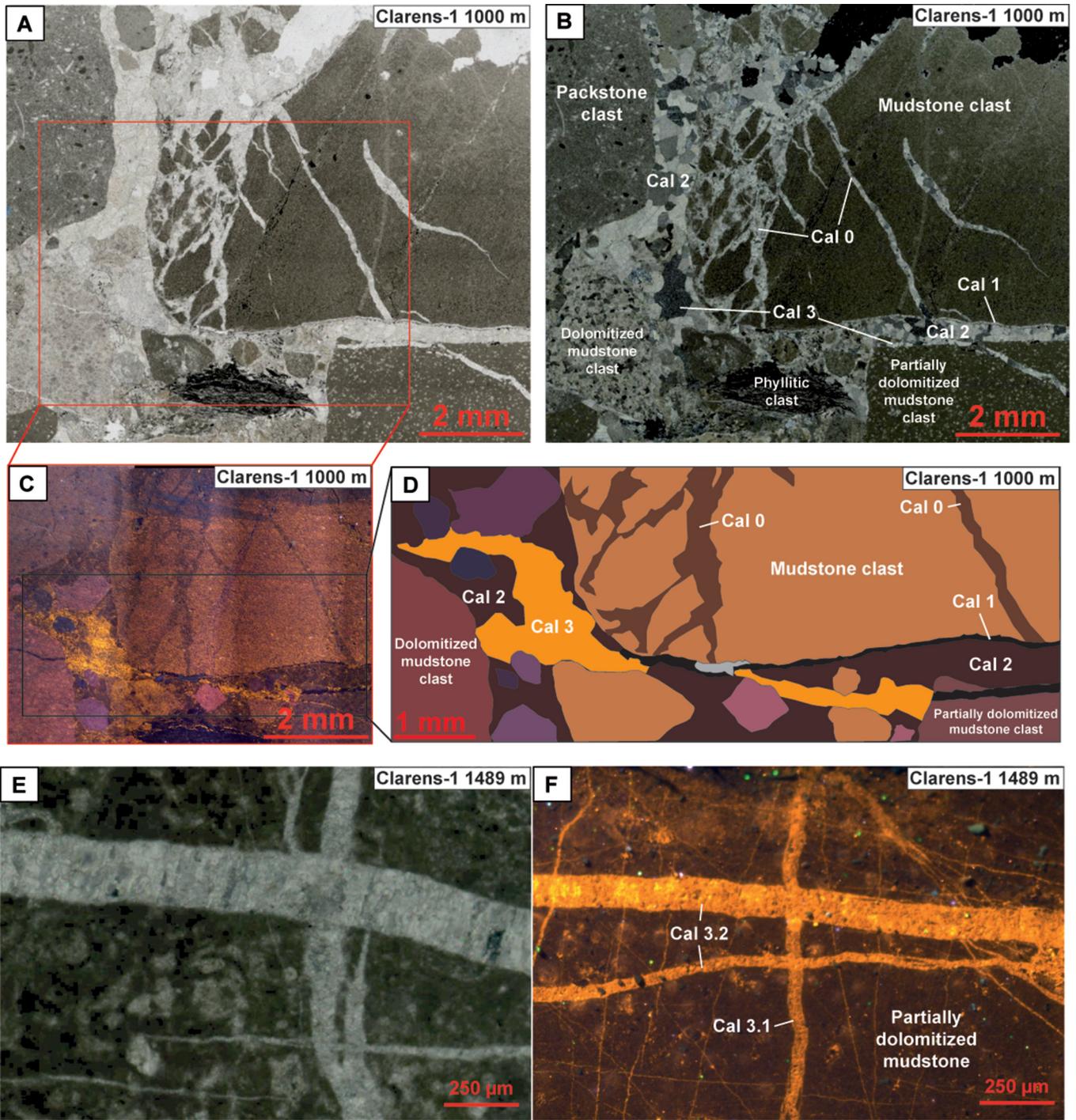


Fig. 5. (A) Plane polarized light image of the breccia observed at 1000 m, and (B) the associated cross-polarized light image with the interpretation of the different clasts and veins. (C) Cathodoluminescence image of the same breccia observed in (A) and (B) with its interpretative sketch in (D). (E) Cross-polarized light image of the late calcite veins observed at 1487 m. (F) Cathodoluminescence image of another zone of the same thin section as in (E).

Table 1. Fluid inclusions microthermometric data from Clarens-1 borehole at 1000 m and 1487 m depth.

Samples	Fluid inclusion number	Fluid inclusion host type	T_c (°C)	$T_{m(ice)}$ (°C)	Salinity (wt. % NaCl eq)	T_h (°C)
Clarens-1 1000 m Cenomanian Polygenic breccia	2b	Primary	-44.3	-21.3	23.4	
	17	Primary	-42.8	-19.8	22.4	81.9
	2	Primary	-43.3	-20.2	22.7	118.8
	3	Primary		-18.2	21.3	120.1
	4	Primary		-18.1	21.2	122.2
	5	Primary		-19.2	22.0	133.1
	6	Primary	-39.9	-22.6	24.2	116.9
	7	Primary		-7.8	11.5	109.7
	8	Primary	-37.9	-6.7	10.1	127.7
	9	Primary	-37.9	-18.4	21.4	134.0
	10	Primary	-41.9	-16.9	20.3	114.9
	11	Primary	-45.8	-18.4	21.4	109.0
	12	Primary		-18.2	21.3	66.9
	13	Primary		-18.5	21.5	76.7
	14	Primary		-19.1	21.9	63.1
	15	Primary		-14.9	18.7	110.3
	18	Primary		-19.0	21.9	81.6
	20	Primary		-18.4	21.4	69.8
	23	Primary		-21.3	23.4	81.7
	1	Primary				145.1
	16	Primary				96.6
	19	Primary				83.9
	21	Primary				103.6
	22	Primary				88.2
24	Primary				86.2	
Clarens-1 1487 m Cenomanian Polygenic breccia	1	Primary	-64.5	-23.3	24.6	117.0
	2	Primary		-23.8	25.0	121.5
	3	Primary		-21.9	23.8	137.3
	4	Primary		-15.0	18.8	160.2
	5	Primary		-22.2	24.0	174.7
	6	Primary		-14.4	18.2	72.2
	7	Primary		-19.2	22.0	102.9
	8	Primary		-18.9	21.8	102.0
	9	Primary		-19.0	21.8	105.8
	12	Primary		-22.0	23.8	116.7
	14	Primary		-13.8	17.7	164.7
	15	Primary		-11.6	15.6	150.4
	10	Primary				101.3
	11	Primary				110.5
	13	Primary				96.7
	16	Primary				134.0
	17	Primary				97.7

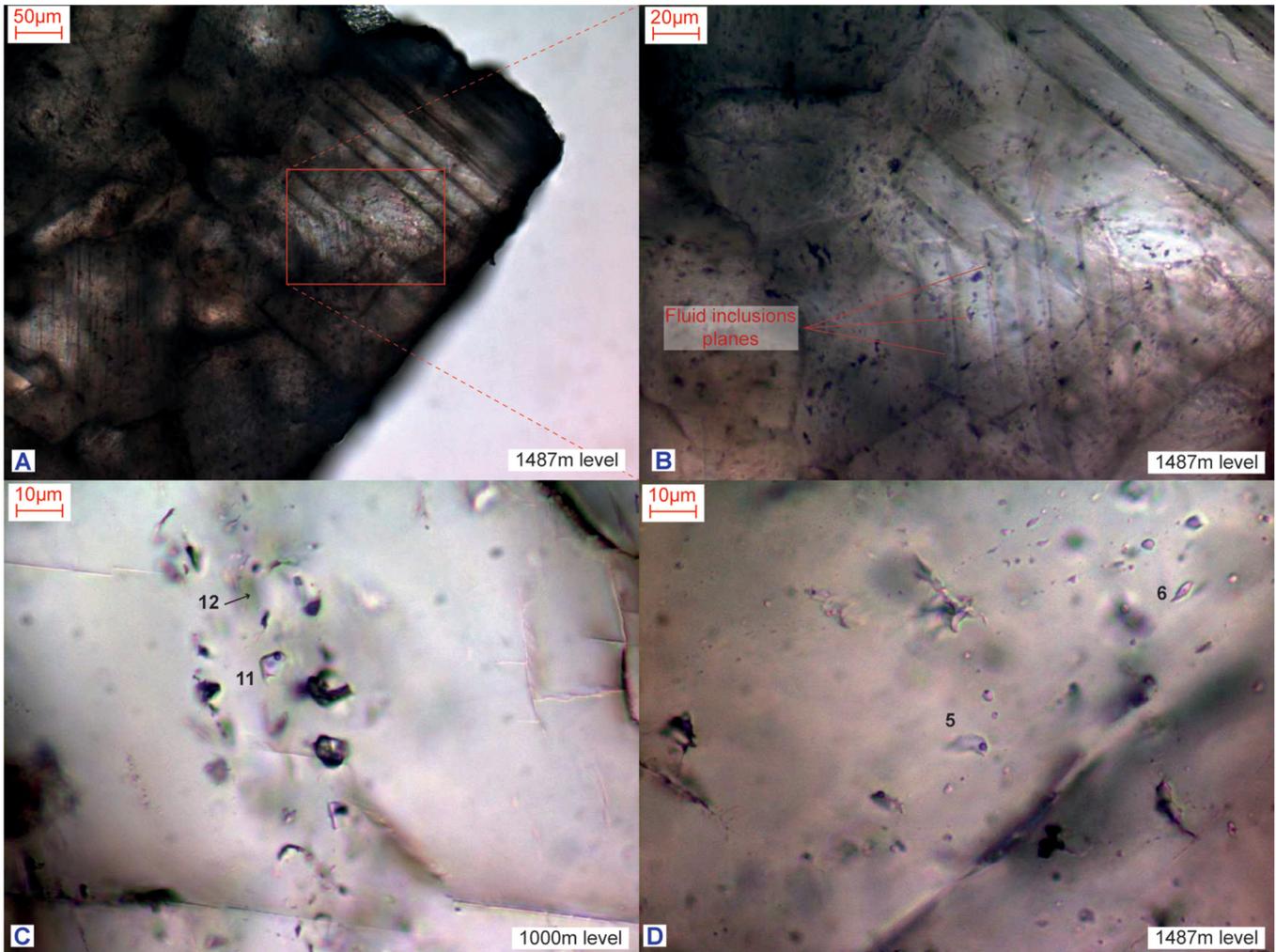


Fig. 6. (A and B) Microphotographs of typical fluid inclusions planes observed in the calcite cleavage. (C and D) Microphotographs of typical fluid inclusions observed in the two levels (1000 m and 1487 m) of the Clarens-1 well.

south (*i.e.*, Baronnies Trough). This is suggested by the fact that pre-rift (Jurassic) supra-salt series are not recognized on the hanging wall of the southern normal fault. Due to decoupling in the Upper Triassic salt and the slope formed by the former normal fault, these series glided to the south and were disconnected from the pre-rift remnants located above the basement high.

6.1.2 Timing of NPFT activity

The transition from post-rift basin evolution to the onset of tectonic inversion along the NPFT can be inferred from the stratigraphic architecture observed in the footwall. Following the passive infilling of post-rift structures by early syn-convergence deposits (Fig. 8A), the wedge geometry defined by overturned Santonian to Maastrichtian beds passing upward into flat-lying strata ahead of the thrust indicate the NPFT was active at this time, which is in agreement with inference from thermochronology (see Fillon *et al.*, 2021 and description below). A part of deformation may have been accommodated

by the weld of the diapiric structure localized along the former normal fault. Minor deformation is accommodated further north in a back-thrust structure which shows a small shift of the syn-convergence successions. The NPFT as well as this back-thrust both result from the inversion of former normal faults. In the Aquitaine Basin, the deformation is also accommodated by the tightening of pre-existing salt diapirs as highlighted by fold geometry of the overlying syn-convergence deposits.

Fillon *et al.* (2021) used apatite fission track and (U-Th)/He on zircon along the Clarens-1 borehole to evaluate the post-orogenic vertical motions of the Aquitaine Basin in response to renewed erosion of the Axial Zone. The thermal modeling of the deeper samples has shown a clear signal of early-orogenic warming followed by slow Cenozoic cooling. While the uppermost samples show monotonic time-Temperature paths since Cenomanian times – close to surface or sub-surface temperatures (Fig. 8B, sample CLS-296) – the lowermost ones recorded heating at the same time. The best-fit scenario resolves a rapid heating event during the Campanian bringing the samples to a temperature of 120 ± 20 °C at 72 ± 2 Ma (Fig. 8B).

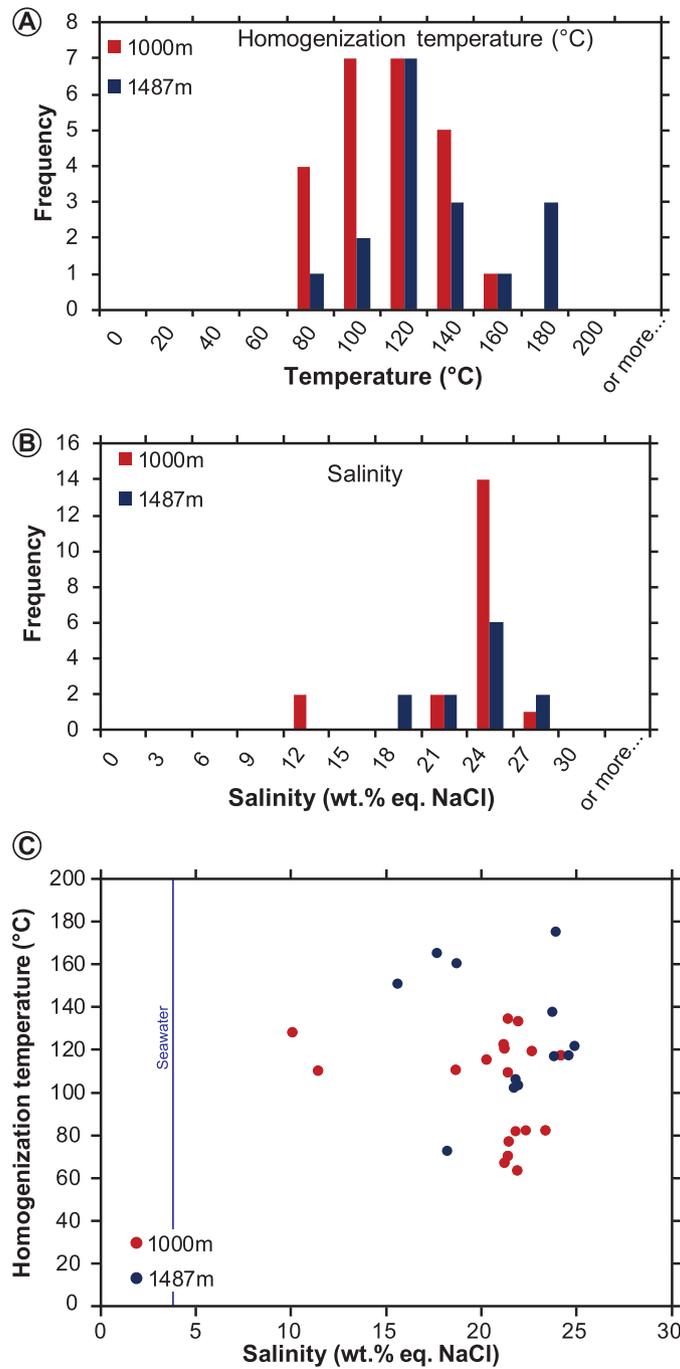


Fig. 7. Microthermometry results of fluid inclusions from Clarens-1 borehole (France). (A) Histogram of homogenization temperatures (T_h). (B) Histogram of salinity (wt.% eq. NaCl). (C) Diagram T_h vs. salinity.

Considering that the uppermost samples do not show any evidence for significant heating after deposition, we interpret the thermal event recorded by the lowermost samples as resulting from the burial induced by overthrusting of the North Pyrenean Frontal Thrust, which is in agreement with vitrinite results. This interpretation is also consistent with seismic reflection data that indicate a first phase of thrusting during the Campanian marked by the wedge-shaped thickening of the Campanian strata (Figs. 3 and 8).

After the early stage of convergence, during the Late Cretaceous, the observation that the Danian (Paleocene) deposits are locally isopachous reveal that the NPFT was inactive at this time (*i.e.*, tectonic quiescence), but salt diapirism was likely ongoing as indicated by sediment thinning above the Recurt-1 diapir. Subsequently, the major syn-tectonic thickening of Eocene sediments figures the reactivation of the NPFT structure, in agreement with the syn-orogenic deformation and exhumation during Eocene-Oligocene times

(Beaumont *et al.*, 2000, Fitzgerald *et al.*, 1999). This second tectonic phase during the Eocene-Oligocene times although supported by a second sediment wedge is not recorded by thermal modelling (Fillon *et al.*, 2021), mainly because the burial may have been lower than 2–3 km, requiring higher sensitivity thermochronometers such as (U-Th)/He on apatite to be resolved. Unfortunately, the low apatite yield of our samples did not allow us to perform such analysis.

6.2 Fluid circulation: nature and timing

6.2.1 Timing of formation of breccias and calcite veins

During the post-rift stage, the paleo-high formed between the Baronnies Trough and the Aquitaine Basin induced a break in slope along the former south-dipping normal fault, which triggered the formation of the sedimentary polygenic breccia observed in the Cenomanian formation with a probable contribution of the diapir (Fig. 8). A first fracturation event corresponding to calcite veins filled by Cal 0 occurred before the formation of the breccias (Fig. 5D). Then, the joints of the breccia were filled by two generations of non-luminescent calcite: Cal 1 and Cal 2. Precipitation of these two generations may have occurred before or during the Campanian activation of the NPFT. Cenomanian to Santonian period is known to be a post-rift period, without any documented major active tectonic structures. In the light of the structural and thermochronology results presented above, a Campanian activation has a much higher probability.

The last generation of calcite (Cal 3), where the fluid inclusions were studied, crosscut the breccias revealing their late injection regarding the first activation of the NPFT. It is worth noting that the Cenomanian marls levels between the breccias are preserved from calcite cements, showing that the breccias were the preferential pathways for fluids circulation. The two generations of fractures observed at 1489 m, suggest two pulses of fluid circulations. But the very similar luminescence and fluid inclusions composition (see in the following sub-section) between the veins at 1489 m and Cal 3 observed at 1000 m, suggest a similar fluid for Cal 3 at these two depths. In addition, the similar salinity of fluid inclusions analyzed at 1000 m and 1487 m strongly suggests a similar origin for these late fluids. Thus, we interpret that the late calcite veins (*i.e.*, Cal 3), cross-cutting the breccias, were generated from the same fluids which originated from the leaching of Triassic evaporites at depth (please see below).

The transition from a stage of joints filling by Cal 1 and Cal 2 to a fracturing of the whole breccia by the Cal 3 veins, necessarily implies a lag time without major tectonic influence. We suggest that the tectonic quiescence of about > 10 Ma between the two activations of the NPFT, corresponds to this transition. This period possibly allowed consolidation of the breccia making it impermeable (such a breccia consolidation was already observed in the South-Pyrenean Zone; Travé *et al.*, 2007). It implies that the calcite veins are late regarding the first NPFT activation and the quiescence time. Thus, they are probably associated with the Eocene-Oligocene reactivation of the thrust, marked by the growth strata observed in Figure 3.

An absolute dating of the calcite veins was attempted using the U/Pb method to validate their age. Unfortunately, the concentrations of uranium and lead were too low (about tens of ppb) in the calcite to determine an absolute age.

6.2.2 Evaporitic origin of the fluid

Fluid inclusions observed in the calcite veins (Cal 3) from both 1000 m and 1487 m levels of the Clarens-1 borehole show very high salinities (up to 25 wt.% eq. NaCl). Thanks to eutectic temperatures, the fluid composition can be determined as a NaCl–CaCl₂–MgCl₂–H₂O (±LiCl) system. Such high salinity and complex composition are in agreement with brines generated by evaporite leaching (*e.g.*, Goldstein, 2001; Yardley and Graham, 2002). No discrimination can be made between the fluid circulating at 1000 m and 1487 m, suggesting a fluid with same origin in the veins observed in the NPFT zone. The only difference between fluid inclusions from the hanging wall and the footwall (Fig. 2) is the temperature. The large variation of T_h with similar salinity and salt composition (*i.e.*, same T_c) cannot be explained through petrographic observations. The different T_h values observed in a same fluid inclusion plane as well as the absence of post-entrapment modification suggest the cooling of a relatively continuous fluid flow from one source which is recorded in a single generation of fluid inclusions. Such an effect has already been observed in the décollement layer within Upper Triassic evaporites of the Alps (Barré *et al.*, 2020), as reflecting the cooling of a single fluid circulation during the retrograde metamorphism of the Alpine collision. The occurrence of Triassic evaporites within the NPFT fault zone is inferred from borehole data and our structural interpretation (Figs. 3 and 8). This pleads for a unique evaporitic origin for the fluids that percolated along the NPFT and that are recorded in the late Cal 3 veins.

The temperature difference inferred from fluid inclusions between the upper (101.8 °C ± 23.4 °C at 1000 m depth) and lower (121.5 °C ± 28.1 °C at 1487 m) levels of the thrust zone, probably reflects temperature difference in the source of the fluids (*i.e.*, Triassic evaporites present in depth). Assuming a thermal gradient around 40 °C/km, we estimate that the fluid was generated between 2.5 and 3.0 km deep. Lannemezan-1 & 2 wells penetrate the Triassic evaporites between 2153 and 2636 m and between 2386 and 4727 m depth, respectively. A Triassic origin for the last fluid circulation generated by the leaching of evaporites is therefore strongly suggested here.

6.2.3 Agreement with the late fluid circulation in the Pyrenees

Most of the fluid circulations studied in the Pyrenees emphasize the contribution of evaporites in the composition of fluids (McCaig *et al.*, 2000; Travé *et al.*, 2000; Salardon *et al.*, 2017; Corre *et al.*, 2018; Crognier *et al.*, 2018; Quesnel *et al.*, 2019; Incerpi *et al.*, 2020). A recent review from Cathelineau *et al.* (2021) showed that evaporitic-related fluids are ubiquitous throughout the geological history of the Pyrenees. These fluids are mixed in different proportions with more dilute water sources (*e.g.*, meteoric water, seawater...) inducing the recording of a relatively wide salinity range in the fluid inclusions. Therefore, by using salinity, it is not possible to determine whether the fluid circulated during the rifting or during the compression phase of the orogen.

To resolve this critical parameter, most of the studies in the Pyrenees have used structural evidence to characterize the timing of fluid circulations they observed (Travé *et al.*, 1998,

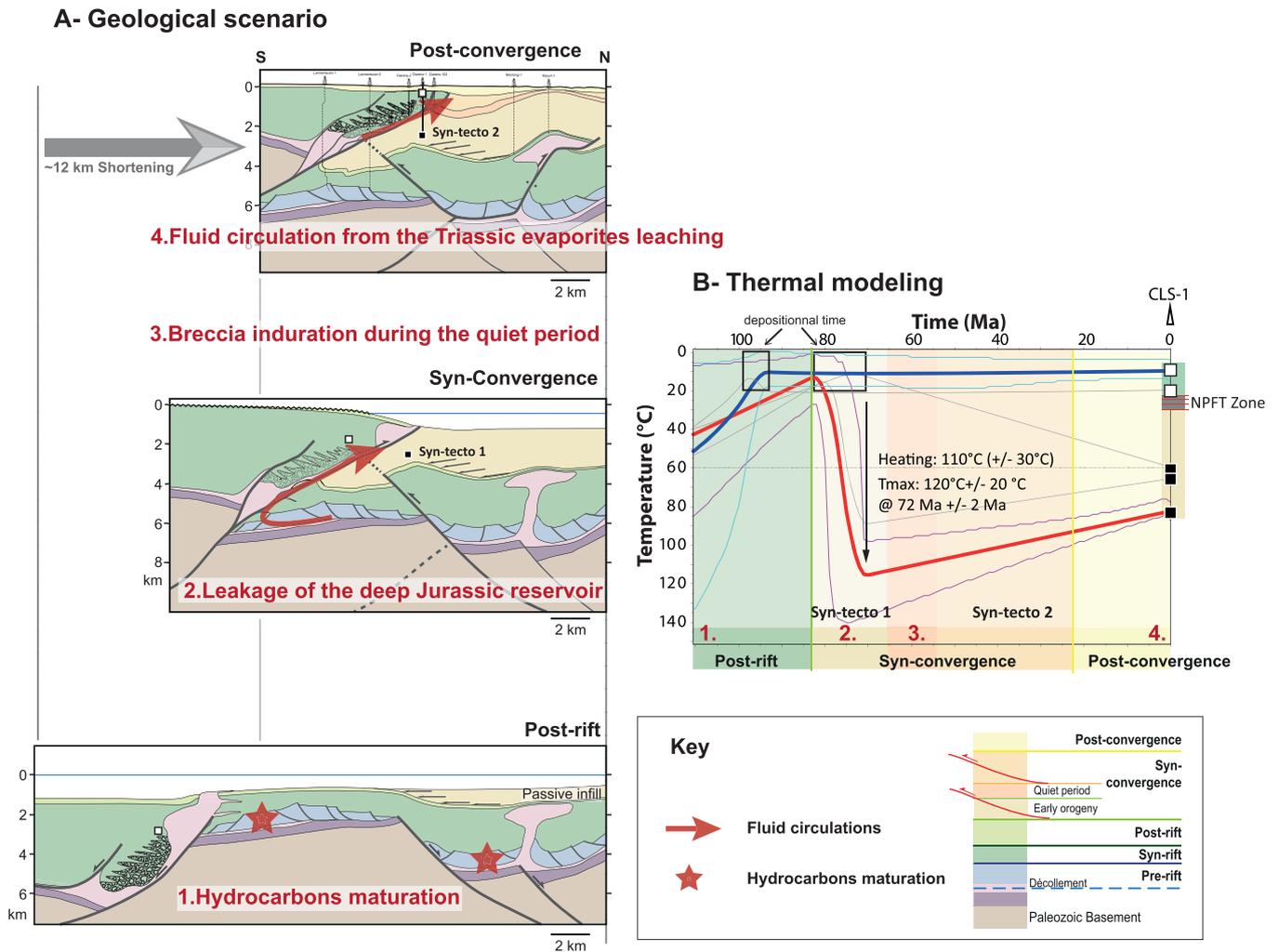


Fig. 8. (A) Restored cross-sections of Figure 3 with associated fluid flows in space and time. (B) Time–temperature paths from the thermal inverse modeling from Fillon *et al.* (2021).

Table 2. Mean vitrinite reflectance (VRo) data from 224 m to 3501 m depth from the Clarens-1 borehole.

Clarens-1				
Depth (m)	Age	VRo (%)	S.D.	n
224.0	Cenomanian	–	–	–
1523.0	Maastrichtian	0.28	0.04	45
2582.4	Campanian	0.38	0.04	60
3048.4	Campanian	0.55	0.06	50
3501.0	Campanian	0.60	0.05	55

S.D.: standard deviation; n: number of measurements.

Table 3. Isotopic and chemical composition of the gas from Lannemezan-2 borehole (data from Lacrampe, 1990).

Well	Age	$\delta^{13}\text{C CO}_2$ (‰V-PDB)	$\delta^{13}\text{C CH}_4$ (‰V-PDB)	N ₂ (%)	CO ₂ (%)	H ₂ S (%)	CH ₄ (%)	C ₂ H ₆ (%)	C ₃ H ₈ (%)	iC ₄ (%)
Lannemezan-2 6137 m	Tithonian	-2.6	-37.1	0.71	4.00	0.69	94.24	0.31	0.03	0.02

2000; Lacroix *et al.*, 2014; Beaudoin *et al.*, 2015; Crognier *et al.*, 2018). Using this technique, most of the fluid circulations observed are associated with the compressive stage of orogeny. A single study shows fluid circulations related to the extensive stage using structural evidence (Salardon *et al.*, 2017). A few studies have used absolute dating to determine the exact timing of fluid circulation (Incerpi *et al.*, 2020; Motte *et al.*, 2021). So far, they have only obtained ages showing fluid circulation during the extension phase. Recent studies used thermochronological tools to estimate the age of circulated fluids in faults (Milesi *et al.*, 2019, 2020). They showed that in the Têt fault in the Axial Zone, fluids circulated during the compression phase. However, no fluid inclusion or isotopic study was coupled with these studies making it impossible to determine the nature of these fluids.

Here, it is the first time in the Pyrenees that a study of fluid inclusions has been associated with seismic and thermochronological data. It allowed to determine that the last fluid (*i.e.*, Paleogene fluids) circulated in the NPFT was highly saline and generated by the leaching of the Triassic evaporites during the Eocene-Oligocene reactivation of the fault. This is consistent with the late fluid circulations observed in the Chaînons Béarnais (NPZ) showing a high salinity around 18 wt % eq. NaCl (Motte *et al.*, 2021). In the Aquitaine Basin, compression fluids have very low salinities around or below seawater (Renard *et al.*, 2019; Elias-Bahnan *et al.*, 2020), while the late fluid circulations observed in the South Pyrenean Zone have low salinity around 7 wt% eq. NaCl (Travé *et al.*, 1998; Crognier *et al.*, 2018). In all cases, meteoric water is assumed to be the main source of fluids with varying influence of evaporite leaching. Here, the Triassic evaporites are present in relatively low depth and are directly incorporated into the NPFT. It can therefore be assumed that, as in the other parts of the Pyrenees, the meteoric waters infiltrated to the evaporites, leached them, then warmed up leading to the late fluid circulation analyzed here in the calcite veins (Cal3). The proximity of the evaporites and the presence of the NPFT as a preferential pathway for fluid circulation (see below), may explain why the fluids are saltier here than in the other late fluid circulations observed elsewhere in the Pyrenees. This is consistent with the conclusions of Cathelineau *et al.* (2021).

6.3 The conduit and barrier role of the NPFT

The association of breccias and veins observed in the NPFT is consistent with a model of conduit-barrier fault where fluid circulations are preferentially located in the fault zone (Caine *et al.*, 1996; Bense and Person, 2006). The breccias formed earlier than the veins. This suggests that the NPFT acted as a barrier and then as a conduit, the latter being evidenced by the late calcite veins analyzed in this study. The late fluids recorded in these veins originate from the Triassic evaporites present deeper (see previous section). However, at depth the Lannemezan-2 borehole reached a deep Jurassic hydrocarbon reservoir at 6137 m, and based on the cross-section (Fig. 3) informed by seismic and well data this reservoir is connected to the NPFT. One may assume that hydrocarbons should then be observed in the late veins. But no

hydrocarbons (oil or gas) are observed in these veins. Moreover, no gas was found in Clarens-1 and Clarens-2 boreholes in contrast to the Lannemezan-2 borehole where gas are present and composed by 94.24% of CH₄ and minor CO₂ (4.00%), N₂ (0.71%), H₂S (0.69), C₂H₆ (0.31%), C₃H₈ (0.03%) and iC₄H₁₀ (0.02%; Tab. 3). We suggest that the upper part of the NPFT acted as a conduit while the lower part of the fault acted as a barrier. Such a variation can be due to the Triassic evaporites present between the two reservoirs (Fig. 3) which may act as an impermeable formation preventing the rise of deep fluids through the NPFT.

Moreover, the gas composition and the isotopic signature of Lannemezan reservoir differ from the global gas composition in the other reservoirs found in the Aquitaine Basin with higher methane content and more enriched carbon isotope signatures (Lacrampe, 1990). According to Milkov and Etiope (2018), the combined carbon isotope signatures of CH₄ ($\delta^{13}\text{C} = -37.1\text{‰}$ V-PDB) and CO₂ ($\delta^{13}\text{C} = -2.6\text{‰}$ V-PDB; Tab. 3) indicate that Lannemezan gas is a late mature thermogenic gas. The other reservoirs found in the Aquitaine Basin show CH₄ and CO₂ carbon isotopes values around $-43.4\text{‰} \pm 1.9$ and $-1.1\text{‰} \pm 3.9$, respectively (Lacrampe, 1990), which are values representative of an oil-associated thermogenic gas (*i.e.*, less mature hydrocarbons; Milkov and Etiope, 2018). The different gas composition in Lannemezan reservoir may result either from its proximity to the NPZ where high rift-related temperatures have been reported (*e.g.*, Dauteuil and Ricou, 1989; Clerc and Lagabrielle, 2014; Clerc *et al.*, 2015; Ducoux *et al.*, 2021b); or reflects an early leakage from the reservoir before sealing either by breccia consolidation and/or by evaporites.

However, the vitrinite reflectance data from the sedimentary pile located below the NPFT footwall, that is, in the same structural position as the Lannemezan reservoir, show a normal geothermal gradient (see above). We therefore suggest that the NPFT first acted as a conduit during the extension and/or early convergence leading to the Lannemezan Jurassic reservoir leak. Then, the long period of time > 10 Ma that lasted between the two tectonic activities of the NPFT probably promoted the consolidation of the breccia causing the NPFT to act as a fluid barrier. Finally, the tectonic inversion led to the incorporation of the Triassic evaporites in the NPFT. The impermeable evaporites present in the fault resulted in the compartmentalization of the fluids between the upper and lower parts of the fault. This explains why there is no gas in the Clarens boreholes while there is still mature gas in the Lannemezan reservoir.

Further studies are certainly needed to validate this hypothesis, but so far it remains the most likely explanation.

7 Conclusions

This is the first study in the Pyrenees that combined fluid inclusions study to thermochronological studies in a structural framework. It documents late Eocene-Oligocene fluid circulations and long-term thermicity during two phases of thrusting across the North Pyrenean Frontal Thrust. We draw the following conclusions:

- By coupling structural study and inverse modeling of thermochronological data we conclude that the NPFT

activity was initiated by the inversion of a normal fault during the Campanian time and reactivated in the Eocene-Oligocene time.

- The Eocene-Oligocene fluid/rock interaction history reveals circulation of a fluid derived from Triassic evaporitic levels which filled late calcite veins (named Cal 3 here) that affect the Albo-Cenomanian sedimentary breccias. This is in agreement with the recent review on the influence of the evaporites on fluid circulation in the Pyrenees (Cathelineau *et al.*, 2021).
- These data are completed by the isotopic signature and composition of the gas from the Lannemezan reservoir. These highlight the role of the NPFT as a conduit and barrier for fluids between its different activations, separated by a time lapse of more than 10 Ma.

We show that during their polyphase tectonic history, large-scale structures at the orogenic front like NPFT can act as fluid flow conduits and lead to the leakage of oil and gas fields. Later in their history they can act as a fluid barrier. Although more studies are needed to better characterize the timing of the leakage role of the NPFT, the multi-method study of these main structures can provide significant inputs for both oil and mining exploration.

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